



## Original paper

# Modification of a spatially referenced crop model to simulate the effect of spatial pattern of subsoil salinity

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## ABSTRACT

High levels of subsoil salinity limit the growth and yield of dryland cereals in the Victorian southern Mallee, Australia. Currently available crop simulation models of wheat production perform poorly in this region, presumably due to their inability to account for subsoil limitations, mainly salinity. The objective of this work was to modify a spatially referenced Water and Nitrogen Management Model (WNMM) to account for the spatial pattern of subsoil salinity, by adjusting crop water uptake, in order to explain the spatial variation in wheat yield in this area. Measurements of above-ground biomass and yield of wheat, and the profile of soil salinity (0–80 cm) were made at 40 locations across an 88 ha paddock (35.78°S, 142.98°E) in the Victorian southern Mallee. The S-shaped water stress response function for crop water uptake proposed by van Genuchten (1987) was explored to modify the WNMM by adjusting the water uptake due to salinity, which significantly improved yield simulation over the original WNMM. The improvement in the model's ability to simulate wheat yield indicates that the subsoil salinity limits crop performance in the area. The incorporation of a salinity function in spatial crop models offers potential for simulating yield across a landscape and thus practicing precision agriculture provided salinity impact is considered dynamically.

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## 1. Introduction

Salinity is one of the most important environmental stresses influencing the productivity of agricultural systems around the world (McWilliam, 1986). It negatively affects crop growth and yield through the development of negative osmotic potential in the root zone (Tedeschi and Menenti, 2002) as well as causing ion toxicity (Bresler et al., 1982). The presence of salts in aqueous solutions decreases osmotic potential and hinders the extraction and uptake of water by plant roots. Consequently plants are unable to meet their evapotranspiration demand despite sufficient soil water levels (Maas and Hoffman, 1977; Maas, 1986; Somani, 1991; Katerji et al., 1997; Lamsal et al., 1999). Apart from variation in extreme temperatures (Dixit and Chen, 2010a), spatially variable wheat yield in the Victorian southern Mallee, Australia could be related to subsoil constraints, notably soil salinity (Anon, 2004). A field survey (Nuttall et al., 2001a,b) and an investigation of cropping systems on Calcarosols in the Victorian southern Mallee (Nuttall et al., 2003b) showed that these soils contain variable but usually high

levels of salt. Salinity in many soils in this region is also reported by Incerti and O'Leary (1990) and Holloway and Alston (1992). Studies conducted by Sadras et al. (2002), Nuttall et al. (2003a) and McDonald (2006) on similar soils and climate to that used in this study have established that a significant proportion of spatial variability in grain yield of dryland crops could be attributed to the subsoil constraints mainly salinity and its effect on water uptake by the crop.

Salinity problem becomes more severe in arid and semi arid regions compared to humid and sub-humid regions, where rainfall is sufficient to leach out accumulated salts (Lamsal et al., 1999). The southern Mallee region of Victoria receives average seasonal rainfall of less than 250 mm (Sadras et al., 2003). The seasonal rainfall at the site used in this study ranged from 220 mm in 2005 to only 89 mm in 2006 and the long-term average for the Birchip Bureau of Meteorology site (1891–2006) about 20 km from the study site is 239 mm (Armstrong et al., 2009). Due to the low growing season rainfall and high evaporation demand, the salinity problem is severe in this region. Incerti and O'Leary (1990) reported that in 2 years the growing season rainfall was 278 and 201 mm and the evaporation demand ( $E_{pan}$ ) was 683 and 716 mm, respectively. Average actual evapotranspiration (water use) for a wheat crop across 14 locations in Victorian southern Mallee ranged from 150 to 350 mm (Nuttall et al., 2003a) and from 197 to 304 mm near the study paddock (Dixit and Chen, 2010b). There is no irrigation used

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in the area and capillary rise is not an issue due to significant depth of the water table (30–50 m) (O'Connell et al., 1995).

In dryland regions with annual rainfall between 250 and 600 mm, saline subsoils having electrical conductivity of saturation extract ( $EC_e$ ) values between 2 and 16 dS/m can dramatically affect crop production through osmotic effects during dry periods (Rengasamy, 2002). This osmotic effect contributes to a reduced growth rate, changes in leaf colour and developmental characteristics, such as reduced root, shoot and leaf growth and hastens maturity (Steppuhn et al., 2005). Sadras et al. (2002) reported that high soil salinity in the deeper subsoil restricts wheat crop growth by reducing the osmotic potential and adsorption of water, reducing grain yield in the Victorian Mallee. The adverse effect of salinity on crop growth and yield has also been reported by several researchers (Ayers and Wescot, 1976; Maas and Hoffman, 1977; Munns and Rawson, 1999; Sadras et al., 2002; Poustini and Siosemardeh, 2004; Saqib et al., 2004).

The simplistic models based on generally accepted understanding of mechanism often explain too little of the total yield variation to be useful (Cook and Bramley, 2001). This warrants a need to improve the realism of models by adding more and more components to reflect the complexity of biological systems (Booltink et al., 2001). Despite the apparent significance of subsoil properties to crop production in Victorian southern Mallee, few quantitative data are available to define the relationships (Nuttall et al., 2003a). The low seasonal rainfall and the presence of salinity at root zone in the dryland cropping belt of Victorian southern Mallee, demand careful management of scarce water resource and improved simulation models to explain grain yield variation.

The use of simulation modelling to improve crop management decisions, optimise cropping systems and quantify environmental risks has proven to be important and valuable (Meinke and Hochman, 2000). However, subsoil limitations such as salinity or sodicity have so far limited the application of simulation models in the main cereal-growing areas of southern Mallee/North Western Victoria (Rodriguez and Nuttall, 2003).

Most of the commonly available functional models do not simulate the effects of salinity on plant growth (Adiku et al., 2001). However, there have been efforts to account for salinity. Cardon and Letey (1992) found that the S-shaped stress function proposed by van Genuchten (1987) which relates water uptake to the average salt concentration in the root zone with the use of a soil salinity stress function agreed with the measured data. Karlberg et al. (2006) also found similar results in their study when they used a physically based transient ecosystems model which uses the van Genuchten's soil salinity reduction function to determine the effect of increased soil salinity on the plant.

Attempts have been made to modify the single point models to take salinity into account (Lamsal et al., 1999; Homae et al., 2002; Karlberg et al., 2006) in a very general sense. However, no attempts have been made to take the salinity into account, in the low rainfall areas of the wheat belt of Australia with subsoil constraints, to explain variation in the grain yield which could result from the low water extraction by the plants because of subsoil salinity. Additionally, no literature is available on the modification of the model by trying to take salinity effect dynamically i.e., by modifying the concentration of salinity in soil solution at a daily time step. Hence this work presents a novel approach and could be a valuable contribution in the way researchers have been trying to modify the crop models to account for the effect of salinity (or any other soil constraint) on grain yield especially in arid and semi arid areas.

The objective of this work is to modify a spatially referenced Water and Nitrogen Management Model (WNMM) (Li et al., 2007) to account for the effect of salinity on water uptake by the crop which can spatially simulate the effect of subsoil salinity on wheat yield.

## 2. Materials and methods

### 2.1. Field experiment

The study was conducted in an 88 ha paddock (35.78°S, 142.98°E), 20 km north of Birchip in the southern Mallee of Victoria, Australia in the year 2004. During the crop growing period, the average maximum and minimum temperatures were 18.4°C and 5.2°C, the average daily solar radiation was 14.2 MJ/m<sup>2</sup> and the total rainfall at the paddock was 206 mm.

The historical grain yields (from harvester) and biomass (from satellite imagery based on the normalized difference vegetation index, NDVI) during 1996–2002 were collected to define zones of yield variability and thus the yield-based management zones within the paddock were established (Abuzar et al., 2004; Fisher et al., 2009). The paddock area was then assigned into three yield classes (low, medium, and high) and each classified into two seasonal variability zones (variable and stable).

A large agronomic experiment was planned for a precision agriculture project and the required crop data for this study was collected from that experiment, along with the soil salinity data. The agronomic design of the experiment was tailored to meet several other objectives of the project (see Armstrong et al., 2009; Fisher et al., 2009; Rab et al., 2009; Robinson et al., 2009). The summary of the experiment is given below while the detailed layout of the experiment can be found in Rab et al. (2006).

Three experimental blocks were formed within the paddock. The first block was approximately 117 m wide and 700 m long, the second block was 117 m wide and 665 m long and the third block was approximately 117 m wide and 430 m long. Ten zones of different yield potential (two replications of stable high and low yield and three replications of variable high and low yield) were marked among those three blocks based on defined yield zones (as explained above). Each block was longitudinally divided into two sections of sowing rates (30 kg/ha and 80 kg/ha) and each section was further divided into four longitudinal strips where different fertiliser treatments were applied. A wheat crop (variety, Yitpi) was sown at a depth of 5 cm, and with a row spacing of 18 cm. After tillering, nitrogen fertiliser was applied by top dressing in the form of urea that contained approximately 46% nitrogen (N). The four different rates of fertiliser used were 0, 30, 56 and 109 kg/ha (see Rab et al., 2006).

### 2.2. Data collection

Because of low rainfall, the effect of different nitrogen rates was not detected in the field so only the lowest and highest nitrogen treatments i.e., 0 and 109 kg/ha were selected for analysis and plant samples were taken from 10 zones × 2 sowing rates × 2 N treatments, making a total of 40 sample plots. At anthesis, above-ground plant samples were manually collected from 1-m linear distance (2 × 50 cm) from each plot and were oven dried at 70°C for 3 days to obtain dry weight of the above-ground biomass.

Soil cores from the middle of these 40 sample plots (four from each of 10 zones) were taken up to a depth of 80 cm to represent the inherent variability in each yield zone. These cores were subsequently segmented into four sections corresponding to the following depths, 0–10, 10–20, 20–40 and 40–80 cm for the measurement of soil salinity and for particle size analysis to determine clay, sand and silt composition of the soil. These soil data were used to obtain several soil parameters used as input into the model e.g., field capacity, permanent wilting point (lower limit), saturation point and air dry moisture content.

The number and the locations of these representative soil sampling sites were chosen on the basis of variability in grain yield and apparent electrical conductivity ( $EC_a$ ) obtained from an elec-

tromagnetic induction (EM38) survey as reported by Rampant and Abuzar (2004). The soil salinity is correlated with the apparent soil electrical conductivity and is highly related to crop productivity (Freeland, 1989; Rhoades et al., 1989, 1999; Sudduth et al., 1999; Heap, 2002; Whelan and McBratney, 2003). Thereby, the varying salinity levels are inherently represented by different yield zones. In a research to observe interactions between soil properties and water use to explain yield variability, across 3 years of data collection, Armstrong et al. (2009) took 37 monitoring points in 1 year and 40 points in the other 2 years representing the three yield (low, moderate, and high) and two stability (stable and variable) zones in the same research paddock where the study for this paper was conducted.

Further, to establish the adequacy of the measured data in terms of their statistical representation of the spatial variability of sub-soil salinity from the 40 representative sample points, a one-way analysis of variance (ANOVA) was carried out. The ANOVA (Genstat v.11) was performed to test whether the levels of salinity differed across the different zones of yield potential for all the soil layers. The salinity values at four different points within a zone, for a particular layer were considered blocks in the ANOVA. The results showed that the soil salinity was significantly different across the yield zones ( $p < 0.001$  for all layers) which establishes that the collected soil data adequately represent the spatial heterogeneity. The maximum, minimum and mean values of EC for each layer across all the yield zones are given in Table 1.

All the soil samples were dried at 40 °C for 3 days and then ground and sieved to pass through a 2 mm sieve. Altogether, there were 160 samples from 40 sample points and at four different depths. Soil salinity in terms of soil electrical conductivity (EC, dS/m, in 1:5 soil:water solution) was measured according to the standard method described by Rayment and Higginson (1992). The particle size analysis (PSA) of soil samples was carried out by the Nutrient Advantage Lab of Incitec Pivot Limited, Werribee, Victoria as described by Day (1965).

Plant samples were hand harvested from each of the 40 plots at maturity. Above-ground samples were collected from four quadrats within each plot. From each quadrat, two adjacent rows, each 1-m long, were selected and then the bulk samples from all four quadrats, eight rows, were placed into one bag. These samples were oven dried at 70 °C for 3 days and weighed to calculate the above-ground biomass. If the biomass of the harvested sample was less than that of at anthesis, due to decay of leaves, the biomass at anthesis was considered for the analysis. Samples were manually threshed with 100% recovery of grains to obtain grain yield. Harvest index was calculated by dividing the dry grain yield by total above-ground biomass for each harvest sample.

### 2.3. Simulation with the original WNMM

To evaluate the performance of the original WNMM (Li et al., 2007), grain yield and biomass simulations were done using input data from all the 40 points. The observed yield and biomass for different sowing rates were found similar and hence their effect was ignored for all the simulation analyses. Several important crop input parameters e.g., energy efficiency to convert radiation energy into biomass, maximum leaf area index and harvest index were adjusted to calibrate the model.

### 2.4. Modification of the WNMM

The water uptake/transpiration approach is the most commonly used method for estimating salinity stress. This approach assumes that reduced crop growth is caused by a reduced transpiration as a result of increased soil salinity (Karlberg et al., 2006).

**Table 1**  
Salinity (EC, dS/m) values for the different soil layers in each zone.

Zone	EC (dS/m)			
	Layer (cm)	Maximum	Minimum	Mean
1	0–10	0.075	0.043	0.058 <sup>a</sup>
	10–20	0.119	0.069	0.091 <sup>b</sup>
	20–40	0.294	0.145	0.198 <sup>c</sup>
	40–80	0.492	0.215	0.355 <sup>d</sup>
2	0–10	0.166	0.080	0.121 <sup>a</sup>
	10–20	0.168	0.108	0.147 <sup>b</sup>
	20–40	0.210	0.124	0.179 <sup>c</sup>
	40–80	0.492	0.206	0.373 <sup>d</sup>
3	0–10	0.146	0.107	0.126 <sup>a</sup>
	10–20	0.217	0.182	0.202 <sup>b</sup>
	20–40	0.460	0.295	0.380 <sup>c</sup>
	40–80	0.878	0.510	0.700 <sup>d</sup>
4	0–10	0.167	0.102	0.130 <sup>a</sup>
	10–20	0.298	0.167	0.238 <sup>b</sup>
	20–40	0.562	0.258	0.421 <sup>c</sup>
	40–80	0.880	0.492	0.742 <sup>d</sup>
5	0–10	0.103	0.046	0.078 <sup>a</sup>
	10–20	0.246	0.041	0.162 <sup>b</sup>
	20–40	0.390	0.181	0.277 <sup>c</sup>
	40–80	0.629	0.388	0.492 <sup>d</sup>
6	0–10	0.189	0.061	0.127 <sup>a</sup>
	10–20	0.182	0.120	0.147 <sup>b</sup>
	20–40	0.261	0.116	0.187 <sup>c</sup>
	40–80	0.543	0.242	0.404 <sup>d</sup>
7	0–10	0.120	0.070	0.089 <sup>a</sup>
	10–20	0.246	0.112	0.162 <sup>b</sup>
	20–40	0.440	0.203	0.325 <sup>c</sup>
	40–80	0.724	0.500	0.620 <sup>d</sup>
8	0–10	0.227	0.123	0.196 <sup>a</sup>
	10–20	0.352	0.226	0.288 <sup>b</sup>
	20–40	0.675	0.442	0.569 <sup>c</sup>
	40–80	1.062	0.768	0.890 <sup>d</sup>
9	0–10	0.166	0.070	0.125 <sup>a</sup>
	10–20	0.196	0.121	0.156 <sup>b</sup>
	20–40	0.331	0.161	0.254 <sup>c</sup>
	40–80	0.645	0.468	0.561 <sup>d</sup>
10	0–10	0.275	0.193	0.242 <sup>a</sup>
	10–20	0.401	0.384	0.391 <sup>b</sup>
	20–40	0.794	0.559	0.720 <sup>c</sup>
	40–80	1.212	0.783	0.985 <sup>d</sup>

a, b, c and d indicate that the EC values are significantly different ( $p < 0.001$ ) across 10 zones for 0–10, 10–20, 20–40 and 40–80 cm layers, respectively.

van Genuchten (1987) proposed the following S-shaped water stress response function:

$$\alpha(h) = \frac{1}{1 + (h/h_{50})^p} \tag{1}$$

where the water stress response function  $\alpha(h)$  is a dimensionless function of the soil water pressure head ( $0 \leq \alpha \leq 1$ ),  $h$  is pressure head (cm),  $h_{50}$  represents the pressure head (cm) at which the water extraction is reduced by 50%, and  $p$  is a dimensionless experimental constant.

This approach was explored to take the effect of salinity on water uptake similar to the studies by Steppuhn et al. (2005) and Kiani et al. (2008). The modification was made by using the following salinity stress response function for transpiration:

$$T_m = T_o \times \left( \frac{1}{1 + (EC/EC_{50})^p} \right) \tag{2}$$

where  $T_m$  is the modified transpiration (mm) due to salinity,  $T_o$  is the original transpiration (mm) calculated by the model in absence of salinity,  $EC$  is the soil salinity (dS/m) and  $EC_{50}$  is the amount of

salinity which causes transpiration to be reduced by 50%. Steppuhn et al. (2005) reported that the measure used for average root-zone salinity in Eq. (2) is mainly electrical conductivity value ( $EC$ ). A value of  $p = 3$  used by van Genuchten and Hoffman (1984), van Genuchten (1987) and van Genuchten and Gupta (1993) was taken in this analysis.

According to the above approach, the transpiration is more affected exceeding a certain threshold limit of salinity ( $EC_{50}$ ) and less affected below that threshold limit. Maas and Hoffman (1977) reported that empirical relationships between salinity and plant growth could be characterized by a threshold salinity value beyond which yields declines. The modification was done by incorporating Eq. (2) in the transpiration module of the model. First the model calculates transpiration based on Eq. (1) depending on the dryness and then it is modified by Eq. (2) to take salinity into account.

Within this approach, two methods were explored. In the first method  $EC$  was taken as a constant value for a particular layer assuming salinity impact to be constant. Whereas, in the second method  $EC$  was made dynamic by varying it according to the water content of a particular soil layer i.e., the concentration of salt in the soil water at a daily time step assuming salinity impact to be dynamic.

To account for the salinity impact, the water balance module of WNMM was modified for each soil layer as follows:

Update available soil water (salinity)

$$SW_a = SW_p - T_m \quad (3)$$

If salinity impact = constant, then

$$EC = SoilEC \quad (4)$$

$$EC_{50} = \gamma \quad (5)$$

If salinity impact = dynamic, then

$$EC = \frac{SoilEC \times SW_p}{SW_a} \quad (6)$$

$$EC_{50} = \delta \quad (7)$$

where  $SW_a$  is the available soil water in a soil layer in mm,  $SW_p$  is the previous day soil water in that soil layer in mm,  $EC$  is the salinity in dS/m,  $SoilEC$  is the  $EC$  value for the soil layer in dS/m,  $EC_{50}$  is the amount of salinity which causes transpiration to be reduced by 50%,  $\gamma$  and  $\delta$  are the  $EC_{50}$  values for constant and dynamic method of modification, respectively. While activating the model, any of the model modification methods can be selected.

First the modified transpiration was calculated from Eq. (2). Thereafter, soil water for a particular day was modified by subtracting the modified transpiration from the soil water at previous day. The  $EC$  values were the input given in to the model and were calculated for constant impact and dynamic impact of salinity by Eqs. (4) and (6), respectively.

## 2.5. Comparison methods of modifications

The two frequently used statistical indicators in the comparison and evaluation of simulation models are the root mean square error (RMSE) and the mean bias error (MBE) (Wilmott et al., 1985; Jacovides and Kontoyiannis, 1995). However, in this study, the coefficient of determination ( $R^2$ ), Normalized root mean square error (NRMSE) and modelling efficiency (EF) values (Loague and Green, 1991) were also used to investigate the quality of simulation results compared to the measured, along with the mean bias error and root mean square error

$$MBE = \frac{\sum_{i=1}^n (Si - Oi)}{n} \quad (8)$$

$$RMSE = \left( \frac{\sum_{i=1}^n (Si - Oi)^2}{n} \right)^{1/2} \quad (9)$$

$$NRMSE = \left( \frac{\sum_{i=1}^n (Si - Oi)^2}{n} \right)^{1/2} \times \left( \frac{100\%}{\bar{O}} \right) \quad (10)$$

$$EF = \frac{\sum_{i=1}^n (Oi - \bar{O})^2 - \sum_{i=1}^n (Si - Oi)^2}{\sum_{i=1}^n (Oi - \bar{O})^2} \quad (11)$$

where  $O$  and  $S$  represent observed and simulated values, respectively, the corresponding over lined characters represent mean observed and simulated values and  $n$  is the number of observations. The MBE test provides information on the long-term performance of a correlation. A low MBE is desired and indicates better performance of the model. A positive value stands for the average amount of over estimation in the simulated value and vice versa (Bashahu, 2003). The RMSE value determines to what extent the simulations over or under-estimate actual measurements. It is a measure of the scatter of the data points around the 1:1 line. Low RMSE indicates little scatter, while high RMSE indicates large scatter (Zhuang et al., 2001). Generally, quantitative models are considered as accurate when they have a low bias error and low RMSE (Roggo et al., 2003). RMSE is an estimate of the inherent error in the simulation; and the normalized RMSE is a measure of error in relation to the mean. The EF value compares the simulated values to the average value of the measurements. A negative EF value indicates that the average value of the measurements gives a better estimate than the simulated values (Xevi et al., 1996). A high value of  $R^2$  and EF indicate better performance, whereas a low value of MBE, RMSE and NRMSE indicate better performance.

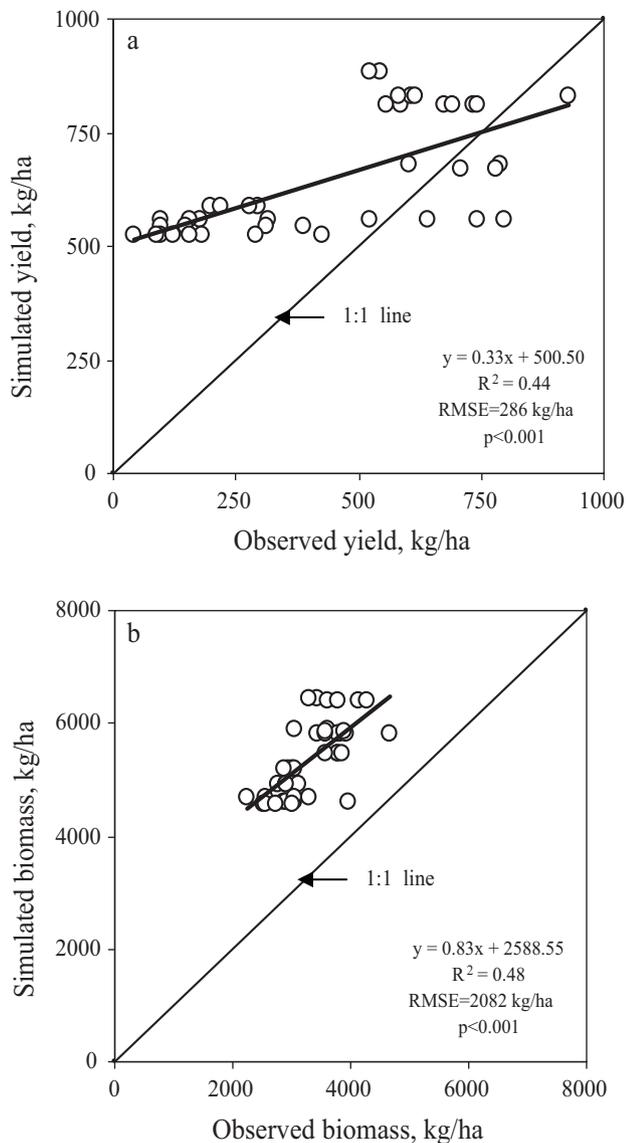
## 3. Results and discussion

Average  $EC$  in the top layer (0–10 cm) across 10 zones was 0.13 dS/m (range 0.04–0.28 dS/m) and in the 10–20 cm layer 0.2 dS/m (range 0.04–0.41 dS/m). Soil salinity and variation in soil salinity was more evident in deeper layers. The 20–40 cm layer had an average  $EC$  of 0.35 dS/m (range 0.12–0.8 dS/m) and the bottom layer of 40–80 cm had the highest average  $EC$  value of 0.6 dS/m (range 0.21–1.2 dS/m).

Average clay and sand in the 0–10 cm layer was 12.9% and 77%, in the 10–20 cm layer 20.4% and 68.4%, in the 20–40 cm layer 30.1% and 58.6% and in the 40–80 cm layer 32.6% and 58.9%, respectively, indicating that while top soil was sandy, subsoil had more clay. Soil salinity increased with increasing clay content which was higher for deeper layers. The association of soil salinity along with other subsoil constraints with clay content is also reported by Armstrong et al. (2009). All the soil water retention parameters except the soil water at saturation i.e., air dry moisture content, field capacity and lower limit increased with increasing levels of salinity. However, the saturation value decreased with increase in salinity.

### 3.1. Simulation with the original WNMM

When the simulated yield and biomass were plotted against their measured values, the original WNMM, without any modification to account for subsoil salinity, produced a high intercept for both yield (Fig. 1a) and biomass (Fig. 1b) and hence was biased. During the calibration process, only the value of intercept kept changing and no significant change in the value of slope was observed, this indicates that the calibration by adjusting crop input parameters alone was not improving the grain yield and biomass simulation. The original model explained about 44% and 48% of the variability ( $n = 40$ ,  $p < 0.001$ ) in grain yield and biomass, respectively, with high biasness. Hence, the coefficient of determination



**Fig. 1.** Relationship between observed and simulated grain yield (a) and biomass (b) by original model.

alone, in this case, cannot be taken as the only measure to explain the performance of the model as it was biased and produced large root mean square error (RMSE) values. Large RMSE indicates poor and unsatisfactory performance of the original WNNM in the study area.

A linear regression without an intercept was conducted, to see if the regression deviates from the ideally expected 1:1 line, and it produced a negative  $R^2$  in the case of grain yield and 0.07 in the case of biomass with a slope value of more than unity indicating over-estimation of yield and biomass. With traditional linear regression formulae, negative  $R^2$ , with no intercept, indicate total inadequacy of the model as reported by Christiansen and Reister (1988).

### 3.2. Simulation with the modified WNNM

#### 3.2.1. Constant impact of salinity

The data from all the 40 sampling points were used to calibrate and assess the performance of the modified WNNM by taking the constant impact of salinity on water uptake by the crop. A  $\gamma$  value of 0.09 dS/m was found to be most appropriate during calibration along with the crop parameters which were

**Table 2a**

Comparison of the different statistical parameters during simulation with original and modified model for grain yield ( $n=40$ ).

Method	$R^{2a}$	MBE (kg/ha)	RMSE (kg/ha)	NRMSE (%)	EF
Original	-3.46	252	286	65.4	-0.3
Constant impact	0.06	282	351	80.3	-0.9
Dynamic impact	0.46	151	188	43.1	0.5

<sup>a</sup>  $R^2$  is for the regression with no intercept.

used to calibrate the original WNNM. When the observed and simulated yield and biomass were plotted, the modified model produced a high intercept (93 kg/ha) for yield showing bias with a low value of slope (0.26). This indicates that the model generally underestimated yield and was less sensitive to variation in simulated yield compared to the measured yield. However, in the case of biomass, the model produced a high negative intercept (-1636 kg/ha) and a slope close to unity (1.04) (figures not presented). This indicates that the model was sensitive to simulation of the variation in biomass with respect to the variation in observed biomass but due to a high negative intercept, it under estimated biomass. Overall the modified model explained about 11% variability in grain yield (RMSE = 351 kg/ha) and 17% variability in biomass (RMSE = 1936 kg/ha) ( $n=40$ ,  $p<0.001$  for both cases). The RMSE value for yield was higher than that of the original model and was similar to the original model for biomass indicating no improvement in the modified model's capabilities to simulate yield and biomass.

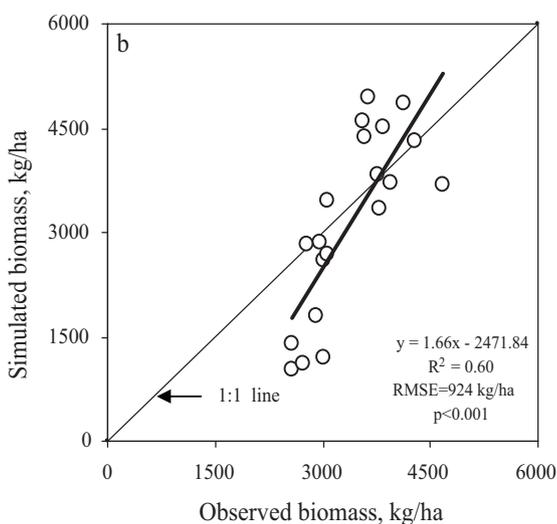
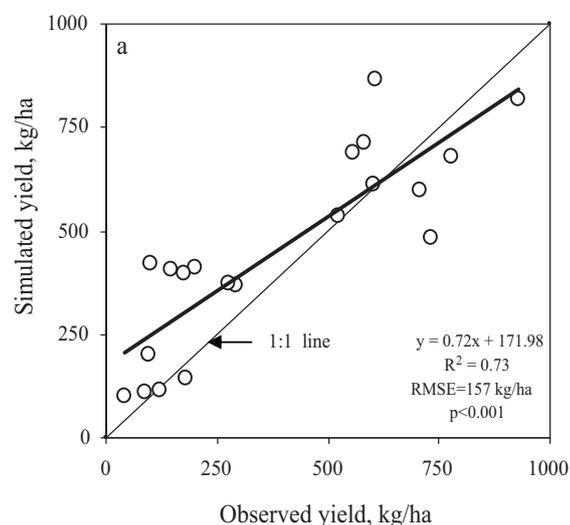
#### 3.2.2. Dynamic impact of salinity

Out of 40 sampled data points, data from 20 randomly chosen points were used for calibrating the model after improving it by taking the dynamic impact of salinity. Data from remaining 20 points were used for validating the modified model. Crop parameters used for calibrating the original WNNM, were also used to calibrate the modified model and were found best fitted. A  $\delta$  value of 0.11 dS/m was found to be the most appropriate in this case during calibration. The calibration was primarily focused on grain yield and biomass predictions.

When the simulated yield and biomass were plotted against their measured values, the modified model (dynamic impact of salinity) explained 73% of the variability in yield (Fig. 2a) and 60% variability in biomass (Fig. 2b) with reduced RMSE in both the cases ( $n=20$ ,  $p<0.001$ ). The reduced RMSE and the improvement in  $R^2$  without intercept for yield (0.54) and biomass (0.49) simulations, compared to the previous method of modification and the original WNNM showed that this method improved simulation significantly. This also indicates that the modified model is good in terms of prediction of change in simulated value with respect to change in observed value (1:1) for both yield and biomass.

During validation, the modified model (dynamic impact of salinity) explained 43% of the variability in yield (Fig. 3a) and 35% variability in biomass (Fig. 3b) ( $n=20$ ,  $p<0.01$ ). The  $R^2$  without intercept for yield (0.38) and biomass (0.30) simulations also show the superiority of this method of modification over the previous one. This also indicates that model is good in predicting change in simulated value with respect to change in observed value for both yield and biomass.

When all the 40 points were considered together by the modified model (dynamic impact of salinity), the RMSE was significantly reduced for both yield (Table 2a) and for biomass (Table 2b), compared to the one obtained from modification by taking constant impact of salinity. The model explained 56% variability in yield and 44% variability in biomass simulations ( $n=40$ ,  $p<0.001$  for both cases). The higher  $R^2$  and lower RMSE indicate better performance of the model by taking the dynamic impact of salinity on water



**Fig. 2.** Relationship between observed and simulated grain yield (a) and biomass (b) during calibration by modified model (dynamic impact of salinity).

**Table 2b**

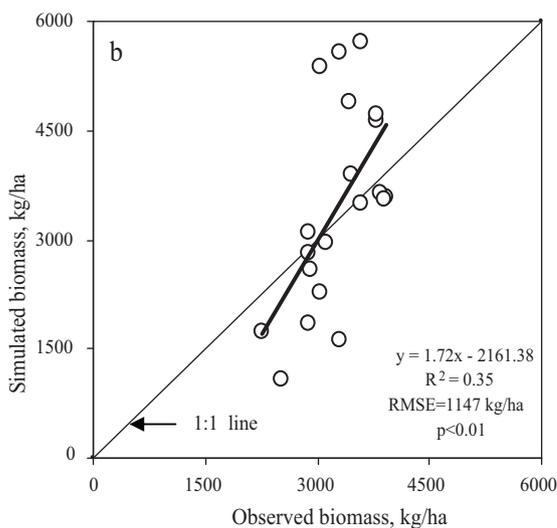
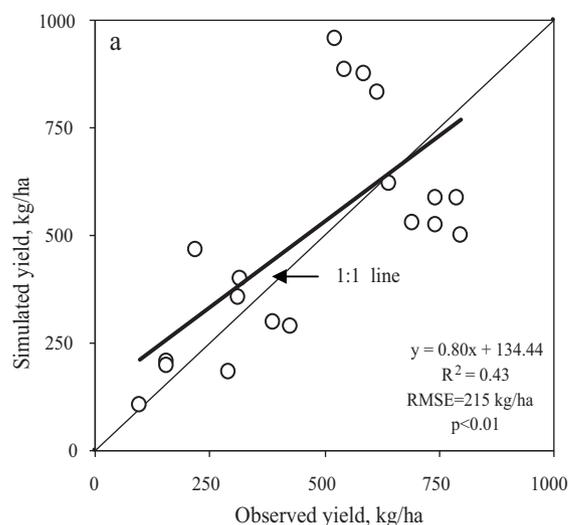
Comparison of the different statistical parameters during simulation with original and modified model for biomass ( $n = 40$ ).

Method	$R^2$ <sup>a</sup>	MBE (kg/ha)	RMSE (kg/ha)	NRMSE (%)	EF
Original	0.07	2028	2082	62.4	-14.2
Constant impact	0.13	1714	1936	58.0	-12.1
Dynamic impact	0.37	814	1042	31.2	-2.8

<sup>a</sup>  $R^2$  is for the regression with no intercept.

uptake than by taking the constant impact. The regression through the origin produced  $R^2$  of 0.46 for yield and 0.37 for biomass simulations and a value of slope almost equal to unity for both grain yield and biomass predictions indicating that the model is good enough in terms of prediction of change in simulated value with respect to change in observed value for both yield and biomass.

The average measured yield and biomass from all the 40 sample plots were 427 and 3335 kg/ha, respectively. The simulated yield and biomass from the original model (without any modification) were 646 and 5363 kg/ha, respectively. When the model was modified by taking the constant impact of salinity, the average simulated yield and biomass were 209 and 1844 kg/ha which were almost half of the measured yield and biomass. This indicates that the salinity impact was considered by the model and both simulated yield and biomass were reduced. But this reduction was not realistic. However, when the model was modified by taking the dynamic



**Fig. 3.** Relationship between observed and simulated grain yield (a) and biomass (b) during validation by modified model (dynamic impact of salinity).

impact of salinity, the average simulated yield and biomass were very close to the measured values, 487 and 3309 kg/ha. This indicates that the reduction in yield and biomass due to salinity was more realistic. Normally, the bottom layers which have more salinity retain more water and the dilution of the salts in the soil solution is less detrimental when the salinity impact is considered dynamic rather than constant. This is the reason that the simulated yield and biomass increased when the impact of salinity on water uptake was considered dynamic opposed to when the impact of salinity was considered constant. On the other hand, when the water from top layers evaporates or infiltrates to the bottom layers, the concentration of salts becomes high in top layers and the plants are more affected when the root zone is shallow which is normally the case during early growth period of the crop. During the early phase of the crop growth, the crop physiological parameters are more affected by the decreased water potential of the rooting solution rather than the yield (Munns et al., 1995). Van Hoorn (1991) reported that in low rainfall areas, due to the evaporation of soil water during germination and emergence, salinity increases strongly in the top layer of the soil and plants are exposed to a higher salinity than during later growth stages. This adversely affects the crop growth at the early stage. However, salinity restricts root growth and makes subsoil water unavailable to crops during the later phase of growth. The low availability of water during grain filling, combined with the inability of the crop to take up water from deep saline subsoil, leads

to reduced grain yield and the spatial variation in subsoil salinity causes spatial variation in grain yield across a paddock (Angus et al., 2001; Robinson et al., 2009).

One of the reasons of success of the van Genuchten's approach is that it modifies the water uptake for transpiration need of the crop due to water stress (which is decreased because of soil salinity) rather than modifying the lower limits of crop available water (which are increased due to salinity rendering less water availability to the plants). Rodriguez and Nuttall (2003) recommended that in heavier soils with substantial chemical constraints, similar to the soils of Wimmera/Mallee region of Victoria, Australia, it may be necessary to account for changes in transpiration efficiency rather than changing the lower limits.

The van Genuchten's approach has been used in many numerical simulation models dealing with root water uptake in saline conditions (Homaee et al., 2002) and found to be successful because it directly modifies the transpiration algorithms within the model and hence provides better estimation of water uptake. Feng et al. (2003) also used the empirical plant water uptake function of van Genuchten (1987) and reported that the simulated yield decreased with increasing irrigation water salinity. They found that the good agreement between the simulated and measured crop yield strongly suggest that the model with van Genuchten's empirical plant water uptake function can be used with confidence under saline conditions. Steppuhn et al. (2005) found that out of six response functions applied to the data from the spring-wheat, the modified-discount, van Genuchten's sigmoidal-shape response function gave the lowest root mean square error and the highest  $R^2$  value.

The improvement in the modified model's ability to simulate wheat yield and biomass, in this study, indicates that the subsoil salinity limits crop performance in the area and its effect must be considered dynamically i.e., salinity should be varied according to its concentration in the soil water for a particular soil layer. The modified WNMM (dynamic impact of salinity) offers potential for accurately simulating yield across a paddock and can help in explaining spatial variation in grain yield at paddock scale. This is a valuable contribution to the grain growers in the salinity affected areas of the arid and semi arid regions.

### 3.3. Statistical performance of the modified model

The yield and biomass simulation performance of the original model and the modified models were compared based on comparative statistics. Tables 2a and 2b give the values of 1:1  $R^2$  and other statistics for grain yield and biomass. The modification by van Genuchten's approach with dynamic impact of salinity performed better than the constant impact of salinity and the original WNMM for both grain yield and biomass simulation. It reduced the RMSE by 35% and 47% for grain yield and about 50% and 47% for biomass and the MBE by 40% and 47% for grain yield and 60% and 53% for biomass compared to the original model and the modification by taking the constant impact of salinity, respectively, indicating a big improvement in model's simulating abilities.

Overall, the modified model (dynamic impact of salinity) produced higher  $R^2$  and EF values, for both yield and biomass, indicating better performance and lower values of RMSE and NRMSE indicating less spread around the 1:1 line and small error in relation to the mean. Also a lower value of MBE indicates that the model is more accurate in simulating yield and biomass and is less biased. A positive value of EF for grain yield suggests that the model provides accurate estimates of average yield. The small negative value of EF indicates that for biomass the performance of the model was not as good as that for the grain yield. However, the EF values for the modified model (dynamic impact of salinity) were reduced by about 4.5 and 5 times compared to the modified model

(constant impact of salinity) and the original model, respectively, indicating about five times better performance.

## 4. Conclusions

Wheat yield simulated with the original WNMM was unrelated to the observed yield across 40 sampling points in the study paddock in the Victorian southern Mallee, Australia. The modified WNMM by incorporating the van Genuchten' stress response function for crop water uptake and considering the dynamic impact of salinity performed well in the study area compared to when the constant impact of salinity was considered. The improvement in the modified model's ability to simulate wheat yield indicates that the subsoil salinity limits crop performance in the Victorian southern Mallee. Also the success of the approach by taking the dynamic impact of salinity on transpiration indicates that in the dry condition of Victorian Mallee it becomes important to modify transpiration module of the model. The incorporation of salinity function in spatial crop models offers potential for simulating yield across a landscape and practicing precision agriculture provided adequate data on soil salinity is used as input and the impact of the salinity is considered dynamically at a daily time step.

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